



UNIVERSITY
OF WOLLONGONG
AUSTRALIA

University of Wollongong
Research Online

Faculty of Science, Medicine and Health - Papers

Faculty of Science, Medicine and Health

2001

Mesospheric temperatures from observations of the hydroxyl (6-2) emission above Davis, Antarctica: A comparison of rotational and Doppler measurements

J L. Innis

Australian Antarctic Division

Frances Phillips

University of Wollongong, francesp@uow.edu.au

G B. Burns

Australian Antarctic Division

P A. Greet

Australian Antarctic Division

W J R French

Australian Antarctic Division

See next page for additional authors

Publication Details

Innis, J. L., Phillips, F. A., Burns, G. B., Greet, P. A., French, W. J. R. & Dyson, P. L. (2001). Mesospheric temperatures from observations of the hydroxyl (6-2) emission above Davis, Antarctica: A comparison of rotational and Doppler measurements. *Annales Geophysicae*, 19 (3), 359-365.

Research Online is the open access institutional repository for the University of Wollongong. For further information contact the UOW Library: research-pubs@uow.edu.au

Mesospheric temperatures from observations of the hydroxyl (6-2) emission above Davis, Antarctica: A comparison of rotational and Doppler measurements

Abstract

We present observations of the hydroxyl (6-2) airglow lines from ~ 87 km altitude obtained at Davis station, Antarctica, in the austral winter of 1999. Nine nights of observations were made of the P-branch near $\lambda 840$ nm with a Czerny-Turner scanning spectrometer (CTS); at the same time, high-resolution Fabry-Perot Spectrometer (FPS) spectra were collected of the $Q_1(1)$ doublet at $\lambda 834$ nm. Rotational temperatures were determined from the CTS observations, while Doppler temperatures were derived from the line-widths of the FPS $Q_1(1)$ spectra. Absolute temperatures determined by these methods are uncertain by ~ 2 and ~ 20 K, respectively. For the comparison we set the value of the reflective finesse of the FPS at $\lambda 834$ nm so the mean FPS temperature from one night of simultaneous data was equal to that from the CTS, and then looked at the measured variations in each data set for the other eight nights. Both instruments show the upper mesosphere temperature to vary in a similar manner to within the observational errors of the measurements, implying an equivalence of the rotational and Doppler temperatures. We believe that this is the first published simultaneous, same-site, comparison of rotational and Doppler temperatures from the OH emission.

Keywords

davis, comparison, mesospheric, doppler, measurements, hydroxyl, 6, 2, emission, above, antarctica, rotational, temperatures, observations, GeoQuest

Disciplines

Medicine and Health Sciences | Social and Behavioral Sciences

Publication Details

Innis, J. L., Phillips, F. A., Burns, G. B., Greet, P. A., French, W. J. R. & Dyson, P. L. (2001). Mesospheric temperatures from observations of the hydroxyl (6-2) emission above Davis, Antarctica: A comparison of rotational and Doppler measurements. *Annales Geophysicae*, 19 (3), 359-365.

Authors

J L. Innis, Frances Phillips, G B. Burns, P A. Greet, W J R French, and P L. Dyson

Mesospheric temperatures from observations of the hydroxyl (6–2) emission above Davis, Antarctica: A comparison of rotational and Doppler measurements

J. L. Innis¹, F. A. Phillips¹, G. B. Burns¹, P. A. Greet¹, W. J. R. French¹, and P. L. Dyson²

¹Australian Antarctic Division, Kingston, Tasmania, 7050, Australia

²Department of Physics, La Trobe University, Bundoora, Victoria, 3083, Australia

Received: 21 September 2000 – Revised: 19 February 2001 – Accepted: 20 February 2001

Abstract. We present observations of the hydroxyl (6–2) airglow lines from ~87 km altitude obtained at Davis station, Antarctica, in the austral winter of 1999. Nine nights of observations were made of the P-branch near $\lambda 840$ nm with a Czerny-Turner scanning spectrometer (CTS); at the same time, high-resolution Fabry-Perot Spectrometer (FPS) spectra were collected of the $Q_1(1)$ doublet at $\lambda 834$ nm. Rotational temperatures were determined from the CTS observations, while Doppler temperatures were derived from the line-widths of the FPS $Q_1(1)$ spectra. Absolute temperatures determined by these methods are uncertain by ~2 and ~20 K, respectively. For the comparison we set the value of the reflective finesse of the FPS at $\lambda 834$ nm so the mean FPS temperature from one night of simultaneous data was equal to that from the CTS, and then looked at the measured variations in each data set for the other eight nights. Both instruments show the upper mesosphere temperature to vary in a similar manner to within the observational errors of the measurements, implying an equivalence of the rotational and Doppler temperatures. We believe that this is the first published simultaneous, same-site, comparison of rotational and Doppler temperatures from the OH emission.

Key words. Atmospheric composition and structure (airglow and aurora; pressure density and temperature; instruments and techniques)

1 Introduction

The hydroxyl airglow emission is often used to derive atmospheric temperatures near mesopause heights (e.g. Takahashi et al., 1974; Armstrong, 1975; Meriwether, 1975; Myrabo, 1986; Offermann and Gerndt, 1990; Sivjee, 1992; Scheer et al., 1994; Bittner et al., 2000). Mesopause temperatures are not only of interest because they represent the coldest part of the atmosphere, but there are indications that tropospheric warming by enhanced greenhouse effects may lead to a cool-

ing of the upper mesosphere (Roble and Dickinson, 1989; Hauchecorne et al., 1991; Golitsyn et al., 1996). The hydroxyl emissions we are concerned with come from a mean height near 87 km, with a thickness of 8 km (Baker and Stair, 1988). The hydroxyl molecule shows a complex rotational-band spectrum, allowing the determination of temperatures by measuring the intensity distribution between various lines in the bands. This approach has been used in a number of studies in both southern and northern hemispheres, and at a range of latitudes (e.g. Scheer and Reisen, 1990; Scheer et al., 1994; Takahashi et al., 1994; Sahai et al., 1996; Mulligan et al., 1996).

Alternatively, a high-resolution Fabry-Perot Spectrometer (FPS) can measure the Doppler width of a line directly (e.g. Jacka, 1985). Temperatures derived from the OH(6–2) $P_1(2)$ doublet were presented by Hernandez et al. (1992). Observations of the OH(6–2) $P_1(3)$ and $Q_1(1)$ emission (erroneously described as $Q_1(2)$ in that paper) are reported by Greet et al. (1994). Instrumental parameters must be well determined for accurate temperatures to be measured (see Greet, 1997).

Both approaches require that the molecules producing the emission are otherwise energetically indistinguishable from the surrounding atmosphere. For valid rotational temperature measurements, the initially unthermalised distribution of rotational quantum upper states in a particular vibration level, $\nu=6$ in our case, need to be effectively thermalised within a vibrational state, prior to emission through collisions and cascading from upper vibrational levels. The effectiveness of thermalisation depends on the lifetime of the upper vibration state, where the higher the state, the shorter the lifetime. For the $\nu=6$ state, a radiative lifetime of the order 6 to 10 ms has been calculated (Turnbull and Lowe, 1989). The collision frequency at 86 km ($\sim 3 \times 10^4$ s⁻¹, US standard atmosphere) indicates the molecules experience around 200 to 300 collisions on average prior to emission; hence, thermalisation is expected (see also Pendleton et al., 1993, for a consideration of non-thermalised hydroxyl emissions). For the Doppler measurements, we require that the velocity distribution of the excited OH molecules be equivalent to that of

the local atmosphere. Hernandez et al. (1992) stated this was the case for the emission from the $P_1(2)$ doublet of the OH(6–2) band. Equipartition of energy (e.g. Sears and Salinger, 1975) supports this contention, given the evidence that the rotational states have thermalised. (We will briefly return to the question of thermalisation of the translational states in our discussion.)

Previous work comparing different techniques for deriving upper mesosphere temperatures includes that of von Zhan et al. (1987), who obtained lidar and OH temperatures on 3 nights, and She and Lowe (1998), who compared a year of OH rotational, lidar, and satellite measurements obtained at three mid-northern locations.

We present observations of the OH(6–2) emissions from Davis station, Antarctica (lat 68.6° S, long 78.0° E) obtained simultaneously with two independent instruments. A scanning Czerny-Turner spectrometer (CTS) recorded the P-branch of the OH(6–2) band near $\lambda 840$ nm; at the same time, the FPS observed the $Q_1(1)$ emission near $\lambda 834$ nm on a number of nights in June 1999. Temperatures were derived from the CTS spectra using the method described in Greet et al. (1998) and French et al. (2000), while the approach of Greet (1997) was used for the FPS data.

Section 2 gives an outline of the two instruments and methods of temperature derivation. Section 3 presents the results of the temperature comparisons, while Sect. 4 discusses the significance of our work.

2 Instruments and reduction methods

2.1 The Czerny-Turner spectrometer

The Davis CTS has been operated for a number of years, and is described in detail by Greet et al. (1998). A summary of the important features is given here.

The spectrometer scans in wavelength by rotating a grating via a stepper motor under computer control. The slit width is set to give a resolution near 0.15 nm at the wavelength of the OH(6–2) emission. The detector is a cooled GaAs photomultiplier. The spectral region near the P_1 branch has been extensively studied (Greet et al., 1998). Contamination from weak auroral lines, other OH emission features, and Fraunhofer absorption lines have been identified and are allowed for or avoided in the spectra used here. Six narrow regions are scanned, corresponding to the $P_1(2)$, $P_1(4)$, and $P_1(5)$ lines, and three background regions (see Fig. 1). Each of the OH lines are doublets, but the CTS resolution is too low to show this. The required spectral intervals were scanned every 7 minutes. Allowance is made for the scan time by linearly interpolating all intensities to a selected time between successive scans, before a rotational temperature is calculated.

The CTS has a 6 degree field of view, and was directed either to the zenith, or at a zenith distance of 70° and geographic azimuth 130° (approximately south-east). The latter was chosen to maximise the signal by looking obliquely through the atmosphere, in an azimuth opposite to usual au-

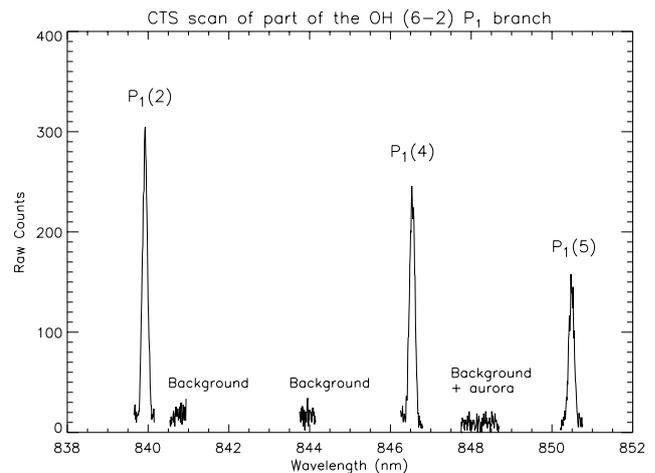


Fig. 1. Spectral scan from the CTS of part of the OH(6–2) P-branch near $\lambda 845$ nm. The spectrometer is stepped between the three lines and background regions during the course of 7 minutes.

roral displays. In this configuration a gold-coated first surface mirror was mounted obliquely on the laboratory roof to allow the off-zenith observations. The mirror showed no significant wavelength-dependent effects over the spectral region of interest.

For derivation of temperatures, we used the method in Greet et al. (1998). The measured P-branch intensities are used to calculate the ratio of molecules in the rotational levels of the upper vibrational state. In order to determine the ratio of the number of molecules in the upper rotational states, transition probabilities are generally required. Turnbull and Lowe (1989) have shown the absolute temperature derived from a range of published transition probabilities increases as the Δv (vibrational quantum number difference) of the band increases. For the OH(6–2) band, Greet et al. (1998) calculate a range of 12 K for three sets of transition probabilities. French et al. (2000) measured the temperature-independent line ratios of the OH(6–2) band and determined the fractions of the upper rotational states that decay via each branch, enabling rotational temperatures to be determined with an accuracy of 2 K. The average temperatures calculated for a selection of spectra using the experimentally derived ratios were within 2 K of the values obtained using the Langhoff et al. (1986) transition probabilities (French et al., 2000). For this current work we adopt the transition probabilities given by Langhoff et al. (1986). In applying this method, care must be exercised in identifying auroral and other contaminating signals, and in removing instrumental effects. Further details are given by Greet et al. (1998) and French et al. (2000).

2.2 The Fabry-Perot spectrometer

The Davis FPS is very similar to the instrument described by Jacka (1985). Slight modifications to allow OH emission observations to be made are described in Greet et al. (1994). The spectrometer uses 150 mm diameter etalon plates, is sep-

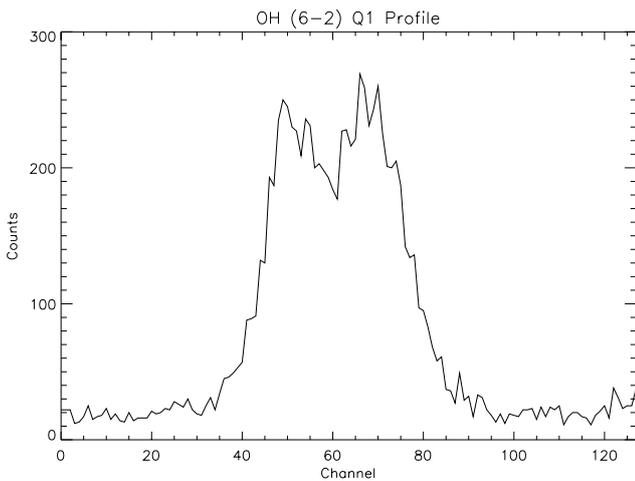


Fig. 2. Sample OH(6–2) Q₁(1) profile at $\lambda 834$ nm, showing the doublet structure. Each channel is approximately 0.3 pm in width in wavelength units.

aration scanned by piezoelectric devices under servo control, and is located in a temperature stabilised cabinet. An aperture mask isolates the centre of the fringe pattern, which is viewed by a GaAs photomultiplier tube. Changing the plate separation scans the instrument in wavelength. Isolation of the spectral region of interest is done by interposing a 0.5 nm interference filter into the light path. In operation, the spectrometer is typically scanned through 128 channels, where each channel is approximately 0.3 pm in width, in ~ 6 seconds. For the OH emission either 80 or 120 such consecutive scans were added to give a single spectrum of the Q₁(1) doublet. A typical spectrum is shown in Fig. 2.

A periscope (0.5° field of view) brings sky light into the instrument. The fields of view of the FPS and CTS were coincident for the data reported here. Differences in observing volumes could lead to observed temperature differences from small-spatial scale atmospheric variations. As our comparison is limited to hourly and nightly averages, and as small-spatial scale effects generally are short-lived, we believe time-averaged data are suitable for comparing the two instruments.

Data reduction involves, among other steps, correcting for photomultiplier dark count, correcting for the finite spectral resolution of the FPS, fitting a pair of Gaussians of identical width to the corrected spectra, and determining the temperature from the Gaussian width. Of specific relevance is the correction for the finite spectral resolution. The method used follows that outlined by Jacka (1985), and discussed further by Greet (1997). On each observing run, several high signal-to-noise ratio observations are obtained of a frequency stabilised He-Ne laser, with its emission at $\lambda 632.8$ nm. The laser is assumed to have zero intrinsic line-width compared to the resolving power of the FPS; hence, the laser profile is taken as a measure of the intrinsic broadening of the instrument. This is termed the instrument function (see Wilksch, 1985, for a detailed discussion of this).

The instrument function will vary with wavelength, as the reflective quality of the etalon plate coatings is also a function of wavelength. Hence, we cannot directly use the $\lambda 632.8$ nm laser profile to correct the $\lambda 834$ nm OH spectra. The quantity that characterises the effect of the reflectivity for an FPS is the reflective finesse, \mathcal{F} . Knowledge of the reflective finesse at the laser and OH wavelengths allows a wavelength-corrected instrument function to be found. Measurement of the reflective finesse, in principle straightforward, is difficult in practice, and has not successfully been done for the Davis FPS at $\lambda 834$ nm. Calculation of the finesse can be done knowing the reflectivity of the etalon coating as a function of wavelength. The manufacturer’s design specification leads to a finesse of 18 at $\lambda 834$ nm. The uncertainty is not known, but could easily be of the order of one unit. A change in the $\mathcal{F}_{\lambda 834}$ finesse by one unit propagates through to a change in the derived temperature by ~ 20 K. For the reasons to be presented below, we adopt a finesse of 19.5 at $\lambda 834$ nm, implying the etalon plates at this wavelength are slightly more reflective than the specifications supplied by the manufacturer. An alternative method of correcting the instrument function from the calibration to observing wavelength was used by Hays et al. (1981) for the Fabry-Perot Interferometer flown on Dynamics Explorer. Their method used detailed knowledge of the working finesse for each instrument channel to produce the required instrument function. We do not have such fine resolution finesse data for our instrument.

With the adopted value of finesse, the instrument function at $\lambda 834$ nm can be found. In principle, this can be deconvolved from the raw data, and the resulting, narrower doublet can then be analysed. Deconvolution is normally carried out in the Fourier domain, where it becomes a simple division. We prefer to avoid the deconvolution, as noise in the data can be magnified to large levels by the division of a small number into a larger one. Hence, we fit to the data a convolution of the instrument function and a pair of Gaussians, as we have found this method to be the more reliable one. The temperature can then be found from the width of the fitted Gaussian. The statistical error in the fit also gives a measure of the uncertainty in the temperature determination.

Temperature comparison observations were scheduled for a two week interval in mid-June 1999. At this time FPS and CTS observations at 70° zenith distance, at azimuth 130° , were obtained for each clear night. Three mostly cloud-free nights of data were obtained: Day of Year (DOY) 169, DOY 173, and DOY 174. Additional OH(6–2) observations carried out in early and late June have also been used in the temperature comparison. On DOY 159 a night of simultaneous CTS/FPS zenith direction data were obtained as part of a study of OH intensity changes during the night hours. On DOYs 153, 175, 176, 178, and 179, the CTS only observed the zenith direction, as part of the usual program, while the FPS performed cardinal point-directed observations, which included zenith measurements once per \sim hourly cycle. For these five nights the temperature comparison was between the (zenith) CTS observations that bracket, or are adjacent

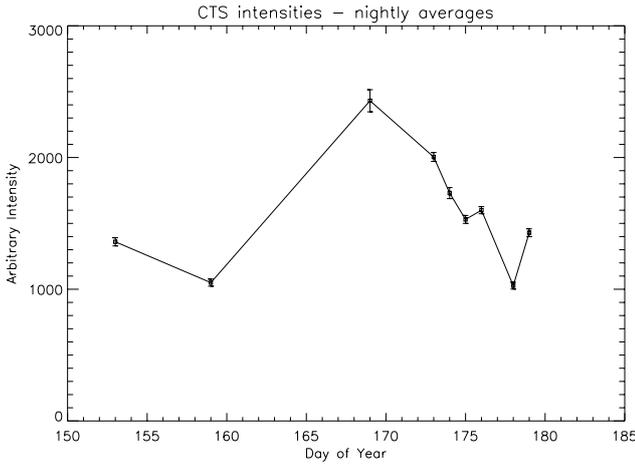


Fig. 3. Mean nightly intensities from the OH(6–2) observations using the CTS, Davis, June 1999. One-sigma standard errors are shown.

to, the FPS zenith measurements; hence, there is one temperature datum every hour. For the nights of DOYs 159, 169, 173 and 174, there are temperature measurements every 6 to 7 minutes, although a small quantity of these data do not pass our acceptance criteria. A few other low quality nights from June 1999 have not been used in this comparison.

3 Results

For both instruments, summing the photon counts in the spectral lines gives a measure of the hydroxyl emission at the time of observation. The two intensity measures are in very good agreement, apart from a multiplicative factor due to differences in the throughput. The hydroxyl emission intensity is known to be variable both during a night, and from night to night. The mean nightly intensities from the CTS are shown in Fig. 3. The greatest signal was on DOY 169.

Temperature data were determined from both CTS and FPS data, as outlined above. For the CTS data, measurements of temperature are obtained from each of the three possible intensity ratios from the $P_1(2)$, $P_1(4)$, and $P_1(5)$ lines. A weighted average is calculated, as described in Greet et al. (1998). The typical standard error in an individual determination is 5 or 6 K. For the FPS, the observational uncertainty for each point is significantly larger, of the order of 20 to 30 K. Inspection of the distribution of the FPS measurements about the nightly mean also reveals the presence of a small number of outlying points, consisting of, at most, $\sim 5\%$ of the data on any night, that do not appear to represent a real physical condition of the atmosphere. To remove these outliers, we reject all FPS data points that are more than 80 K from the mean on a given night.

For the purposes of the comparison between the CTS and FPS temperatures, we have chosen a value of reflective finesse at λ_{834} nm, $\mathcal{F}_{\lambda_{834}}$, of 19.5, so that the mean nightly temperatures for one night of common observations are the

Table 1. Nightly mean temperatures with standard errors from the CTS and FPS. The FPS data have been adjusted to bring the mean temperature from DOY 169 to be equal to that found from the CTS.

Day of Year	FPS: $\bar{T} \pm \text{s.e.}$ (K)	CTS: $\bar{T} \pm \text{s.e.}$ (K)
153	198 \pm 12	203 \pm 2
159	198 \pm 5	206 \pm 1
169	207 \pm 3	207 \pm 1
173	199 \pm 4	201 \pm 1
174	198 \pm 3	195 \pm 1
175	191 \pm 10	211 \pm 2
176	218 \pm 9	212 \pm 2
178	210 \pm 10	206 \pm 2
179	229 \pm 8	223 \pm 2

same. We chose for this night DOY 169, as the mean intensity of the OH signal was the greatest in our data set. We use this finesse value for the other nights. The adjustment made to the finesse value is not substantially different to the manufacturer's specification. The present comparison between the methods is, therefore, between the temperature variations measured by each technique, rather than absolute values.

The nightly mean temperatures we obtained are given in Table 1 and are plotted in Fig. 4, along with standard errors in the means, as found from the data. The standard errors for the CTS and FPS data reflect both statistical and intrinsic variations in the data. The intrinsic atmospheric temperature variations over a night, as seen in the CTS data, are around 10 to 20 K. Statistical errors dominate in the FPS time series. We give standard errors here to indicate how well the nightly-mean FPS temperatures compare with the CTS data. The standard errors for the CTS data are given for completeness.

4 Discussion

From Table 1 and Fig. 4 it is clear that the general trend over the time series is similar for both the CTS and FPS data. The biggest difference is for DOY 175, where the FPS temperature is 20 K less than the CTS result. The error bars shown in Fig. 4 are one-sigma standard errors. On statistical grounds, we would expect that approximately 68% of the time the measured value would be within one-standard error of the 'true' value. We have only a few data points to make the comparison, but excluding DOY 169, six of the eight FPS determined temperatures agree with the CTS results to within this limit, while the others agree to within two standard errors. (For DOY 169 the whole night FPS–CTS temperature difference was set to zero; hence, it is not included in this calculation.) As noted, the sample size is small, but the differences recorded do not appear to deviate from what would be expected, assuming the measurement errors are normally distributed with standard errors determined from the data.

The mean difference, FPS–CTS, and standard error for the eight nightly means (excluding DOY 169) is -2 ± 3 K, which is not significantly different from zero. There may be an in-

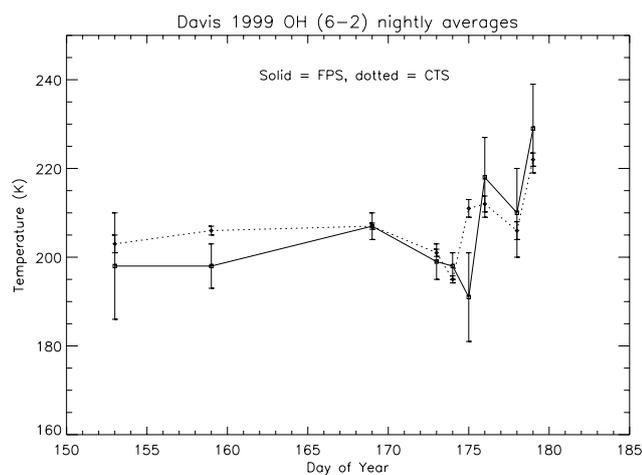


Fig. 4. Mean nightly temperatures from the OH(6–2) observations using the CTS (dotted line) and FPS (solid line), Davis, June 1999. One-sigma standard errors are shown. The absolute values of the FPS temperatures have been adjusted to make the nightly mean for DOY 169 equal to the CTS mean for that night. See the text.

dication of a trend in the differences over the 27 day extent of the simultaneous data, with the FPS initially being low by perhaps 5 K, and then high by the same amount, with respect to the CTS. However, the standard errors in the FPS results at the start and end of the comparison are of the order of 10 K. Additionally, we see no significant difference between the off-zenith (DOYs 169, 173 and 174) and zenith-directed data.

Hourly mean temperatures are obtainable for four days of the campaign: DOY 159, DOY 169, DOY 173, and DOY 174. These data are shown in Fig. 5. The uncertainties are necessarily larger, as fewer data points contribute to the hourly means than to the nightly means, but the two data sets track well. Some 70% of the FPS observations are within one standard error of the corresponding CTS data point (DOY 169 is again excluded in this case). The mean difference FPS–CTS of these hourly data is -2 ± 2 K, again not significantly different from zero.

Temperature measurements from the FPS have a greater observational error than those obtained with the CTS. In part, this is a consequence of the high-spectral resolution of the FPS compared to the CTS, resulting in significantly reduced count rates. Each FPS profile must also have sufficient signal to enable the line shape to be well determined, rather than to give a single measure of the total count, as for the CTS. Many modern FPSs use a fixed etalon spacing and a fringe-viewing area detector to avoid the need to scan in wavelength. Such a system should provide a gain in signal-to-noise ratio per individual measurement, which may lead to a reduction in the observational error per data point. Temperature determinations with such systems appear more difficult than with scanning spectrometers, due to the problem of the analyses of the circular fringe pattern and the doublet nature of the OH emission, but the potential advantages could be significant.

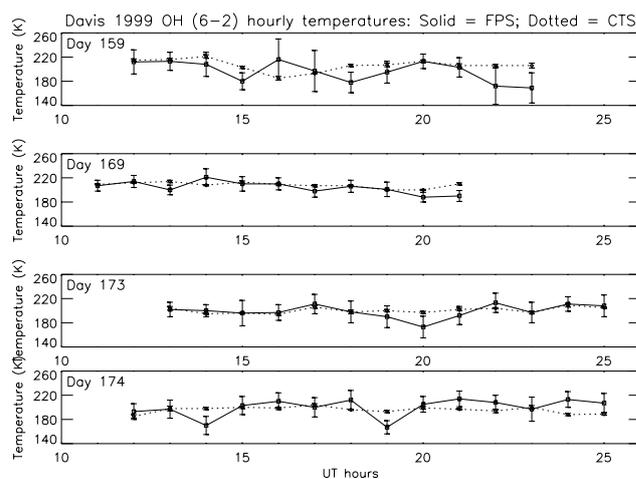


Fig. 5. Mean hourly temperatures from the OH(6–2) observations using the CTS (dotted line) and FPS (solid line), Davis, for four days in June 1999. Top to bottom: DOY 159; DOY 169; DOY 173; DOY 174. One-sigma standard errors are shown.

The agreement between the derived CTS and FPS temperatures over the ~ 25 K range seen during the interval of the observations in Figs. 4 and 5 give a measure of confidence that the measurements are reflecting the real state of the upper mesosphere. We noted earlier that thermalisation of the translational degrees of freedom of the OH molecule would be expected. Our results, showing the agreement between the derived rotational and Doppler temperature changes, lend weight to this assertion.

At a detailed level, however, small differences between the two techniques may be expected. The (approximately) Gaussian height profile of the OH layer is superimposed on a small, and seasonally-variable, vertical temperature gradient. Both measurements are a height-integrated, intensity-weighted, observation of the OH layer. Slightly different weighting-functions contributing to each height-integrated observation may well be expected. For the Doppler method, some slight deviation from pure Gaussian line-shapes could be present, due to contributions from emissions at the lower and upper ranges of the OH layer. However, such effects would be unlikely to be detectable with the signal-to-noise ratios we can currently obtain. Considerations of this type should, however, be noted as potential reasons why exact agreement may not occur.

The absolute temperature measured by the FPS is uncertain because the reflective finesse of the etalon plates at $\lambda 834$ nm has not been well determined. For the present work, however, it is encouraging that good agreement between the two techniques can be found by the assumption of a single reflective finesse value for all the FPS data.

A method of measurement of the finesse \mathcal{F} is described by Jacka (1985). Such measurements have been performed on our FPSs at the Antarctic stations, but not at $\lambda 834$ nm for the Davis instrument. Our experience has shown the properties of the etalon coatings of the Antarctic FPSs are often stable

over a number of years, so that a future determination of \mathcal{F} at $\lambda 834$ nm might be useful for back-calibration of these current data. However, previous finesse measurements at other wavelengths (Murphy, 1996) have shown the difficulty of obtaining a precise value of this quantity; the uncertainty in the result is often ± 0.5 or 1 units. This propagates through the FPS temperature measurements to give a range of the order of 10 to 20 K, making impossible an absolute comparison with the CTS temperatures.

Alternatively, the effect of uncertainties in \mathcal{F} at the observing wavelength can be greatly reduced, if not completely eliminated, by using a calibration source at a very similar wavelength to determine the instrument function. For example, a single-mode, frequency stabilised laser operating, in this case, at $\sim \lambda 834$ nm would give the instrument function directly at the observing wavelength, and should lead to significantly reduced absolute errors in temperature. This should provide an independent check on the adopted transition probabilities and the absolute equivalence of rotational and kinetic temperatures for the OH emission.

5 Conclusions

Wintertime upper mesosphere temperatures over Davis, Antarctica have been derived by two independent instruments and techniques from observations of the hydroxyl (6–2) emissions: a Czerny-Turner spectrometer to find rotational temperatures, and a Fabry-Perot spectrometer to determine Doppler temperatures. An absolute temperature comparison cannot currently be made, due to uncertainties in the reflective finesse, \mathcal{F} , for the FPS at $\lambda 834$ nm. We chose \mathcal{F} to make the mean FPS temperature equal to the CTS temperature for the night of greatest OH emission intensity. The remaining eight nights of data show good agreement in the variations of nightly means. Hourly mean data, available from four nights of the campaign, also show good agreement between the CTS and FPS. We believe this is the first such simultaneous comparison of these techniques of measuring upper mesosphere temperatures. Our results illustrate the essential equivalence of the techniques. Refinement of the value of the reflective finesse at $\lambda 834$ nm for the FPS should allow an absolute temperature comparison to be made.

Acknowledgements. We thank John Everett, formerly with Australian Antarctic Division (AAD), for advice and assistance with the design and manufacture of the gold-surfaced mirror used for the CTS off-zenith observations. Andrew Fleming and Geoff Tonta (AAD Instrumentation Workshop) designed and manufactured the mirror mounting plate. Lloyd Symons (Davis 1999 electronic engineer) maintained and supported the instruments used in this experiment. Tony Morwood (Davis 1999 carpenter) co-ordinated and performed the installation of the gold-surfaced mirror on the laboratory roof in the middle of the Antarctic winter. John Toms (Davis 1999 mechanic) manufactured the roof mounting frame. We thank the other members of the expedition for assistance and support through the year, and Mark Conde of the University of Alaska, Fairbanks (UAF) for useful discussions. This work was supported by the AAD, the Australian Antarctic Science Advisory Committee, and

the Australian Research Council. John Innis acknowledges the hospitality of UAF during part of the time this paper was in preparation. We thank the referees for constructive comments.

Topical Editor D. Murtagh thanks J. Meriwether and another referee for their help in evaluating this paper.

References

- Armstrong, E. B., The influence of a gravity wave on the airglow hydroxyl rotational temperature at night, *J. Atmos. Terr. Phys.*, 37, 1585–1591, 1975.
- Baker, D. J. and Stair, A. T., Rocket measurements of the altitude distributions of the hydroxyl airglow, *Phys. Scripta*, 37, 611–622, 1988.
- Bittner, M., Offermann, D., and Graef, H. H., Mesopause temperature variability above a midlatitude station in Europe, *J. Geophys. Res.*, 105, 2045–2058, 2000.
- French, W. J. R., Burns, G. B., Finlayson, K., Greet, P. A., Lowe, R. P., and Williams, P. F. B., Hydroxyl (6–2) airglow emission intensity ratios for rotational temperature determination, *Ann. Geophysicae*, 18, 1293–1303, 2000.
- Golitsyn, G. S., Semenov, A. I., Shefov, N. N., Fishkova, L. M., Lysenko, E. V., and Perov, S. P., Long-term temperature trends in the middle and upper atmosphere, *Geophys. Res. Lett.*, 23, 1741–1744, 1996.
- Greet, P. A., Innis, J., and Dyson, P. L., High-resolution Fabry-Perot observations of mesospheric OH(6–2) emissions, *Geophys. Res. Lett.*, 21, 1153–1156, 1994.
- Greet, P. A., Mesospheric observations by high-resolution Fabry-Perot spectrometers: calibrations required for climate change studies, *J. Atmos. Sol. Terr. Phys.*, 59, 281–294, 1997.
- Greet, P. A., French, W. J. R., Burns, G. B., Williams, P. F. B., Lowe, R. P., and Finlayson, K., OH(6–2) spectra and rotational temperature measurements at Davis, Antarctica, *Ann. Geophysicae*, 16, 77–89, 1998.
- Hays, P. B., Killeen, T. L., and Kennedy, B. C., The Fabry-Perot interferometer on Dynamics Explorer, *Space Sci. Inst.*, 5, 395–416, 1981.
- Hauchecorne, A., Chanin, M.-L., and Keckhut, P., Climatology and trends of the middle atmosphere temperature (33–87 km) as seen by Rayleigh lidar over the south of France, *J. Geophys. Res.*, 96, 15297–15309, 1991.
- Hernandez, G., Smith, R. W., and Conner, J. F., Neutral wind and temperature in the upper mesosphere above South Pole, Antarctica, *Geophys. Res. Lett.*, 19, 53–56, 1992.
- Jacka, F., Application of Fabry-Perot spectrometers of upper atmosphere temperatures and winds, Middle Atmosphere Program for MAP, R. Vincent (Ed.), ch.13, 19–40, 1985.
- Langhoff, S. R., Werner, H.-J., and Rosmus, P., Theoretical transition probabilities for the OH Meinel system, *J. Mol. Spectr.*, 118, 507–529, 1986.
- Meriwether, J., High-latitude airglow observations of correlated short-term fluctuations in the hydroxyl Meinel 8–3 band intensity and rotational temperature, *Planet. Space Sci.*, 23, 1211–1221, 1975.
- Mulligan, F. J., Horgan, D. F., Galligan, J. G., and Griffin, E. M., Mesopause temperatures and integrated band brightness calculated from airglow OH emissions recorded at Maynooth (53.2° N, 6.4° W) during 1993, *J. Atmos. Terr. Phys.*, 57, 1623–1637, 1995.

- Murphy, D. J., Finesse measurements of the Mawson high-resolution Fabry-Perot spectrometer, ANARE Research Notes, 95, Australian Antarctic Division, Kingston, 221–235, 1996.
- Myrabo, H. K., Winter-season mesopause and lower thermosphere temperatures in the northern polar region, *Planet. Space Sci.*, 34, 1023–1029, 1986.
- Offermann, D. and Gerndt, R., Upper-mesosphere temperatures from OH* emissions, *Adv. Space Res.*, 10(12), 217–221, 1990.
- Pendleton, Jr., W. R., Espy, P. J., and Hammond, M. R., Evidence for non-local-thermodynamic-equilibrium rotation in the OH nightglow, *J. Geophys. Res.*, 98, 11567–11579, 1993.
- Roble, R. G. and Dickinson, R. E., How will changes in carbon dioxide and methane modify the mean structure of the mesosphere and thermosphere? *Geophys. Res. Lett.*, 16, 1441–1444, 1989.
- Sahai, Y., Giers, D. H., Cogger, L. L., Fagundes, P. R., and Garbe, G. P., Solar flux and seasonal variations of the mesopause temperatures at 51° N, *J. Atmos. Terr. Phys.*, 58, 1927–1934, 1996.
- Scheer, J. and Reisen, E. R., Rotational temperatures for OH and O₂ airglow bands measured simultaneously from El Leoncito, *J. Atmos. Terr. Phys.*, 52, 47–57, 1990.
- Scheer, J., Reisen, E. R., Espy, J. P., Bittner, M., Graef, H.-H., Offermann, D., Ammosov, P. P., and Ignatyev, V. M., Large-scale structures in hydroxyl rotational temperatures during DYANA, *J. Atmos. Terr. Phys.*, 56, 1701–1715, 1994.
- Sears, F. W. and Salinger, G. L., *Thermodynamics, Kinetic Theory, and Statistical Thermodynamics*, Addison-Wesley, Reading, Mass., 264–266, 1975.
- She, C. Y. and Lowe, R. P., Seasonal temperature variations in the mesopause region at mid-latitude: comparison of lidar and hydroxyl rotational temperatures using WINDII/UARS OH height profiles, *J. Atmos. Sol. Terr. Phys.*, 60, 1573–1583, 1998.
- Sivjee, G. G., Airglow hydroxyl emissions, *Planet. Space Sci.*, 40, 235–242, 1992.
- Takahashi, H., Clemesha, B. R., and Sahai, Y., Nightglow OH(8–3) band intensities and rotational temperature at 23° S, *Planet. Space Sci.*, 22, 1323–1329, 1974.
- Takahashi, H., Clemesha, B. R., Sahai, Y., and Batista, P. P., Seasonal variations in the mesopause temperature observed at equatorial (4°S) and low (23°N) latitude stations, *Adv. Space Res.*, 14, 197–100, 1994.
- Turnbull, D. N. and Lowe, R. P., New hydroxyl transition probabilities and their importance in airglow studies, *Planet. Space Sci.*, 37, 723–738, 1989.
- von Zahn, U., Fricke, K. H., Gerndt, R., and Blix, T., Mesospheric temperatures and the OH layer height as derived from ground-based lidar and OH* spectrometry, *J. Atmos. Terr. Phys.*, 49, 863–868, 1987.
- Wilksch, P. A., Instrument function of the Fabry-Perot spectrometer, *Appl. Optics*, 24, 1502–1511, 1985.